

Characterizing ozone production and response under different meteorological conditions in Mexico City

W. Lei^{1,2}, M. Zavala^{1,2}, B. de Foy^{1,3}, R. Volkamer^{2,4}, and L. T. Molina^{1,2}

¹Molina Center for Energy and the Environment, CA, USA

²Department of Earth, Atmospheric and Planetary Sciences, Massachusetts Institute of Technology, MA, USA

³Department of Earth and Atmospheric Sciences, Saint Louis University, MO, USA

⁴Department of Chemistry and Biochemistry, University of Colorado at Boulder, CO, USA

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Abstract. Photochemistry in polluted atmospheres, particularly the formation of ozone (O_3), depends not only on pollutant emissions, but also on meteorological conditions. In this study a 3-D chemical transport model CAMx was employed to investigate the O_3 formation and its response to emissions reduction under three distinctively different meteorological conditions in the Mexico City Metropolitan Area during the MCMA-2003 field measurement campaign. The O_3 formation characteristics and sensitivity to emissions change were found to be weakly dependent on the meteorological conditions. The evolution of O_3 formation and its sensitivity to NO_x and VOC levels were also examined along the photochemical plume transport pathway. The midday O_3 production was found to undergo a rapid increase in a narrow range of chemical aging, while downwind plumes were characterized with low and constant O_3 production, and plumes along their transport pathway were characterized by a combination of the two. The O_3 formation was more VOC sensitive near the source area, but as the plume became chemically aged, O_3 formation became progressively VOC insensitive and more NO_x sensitive.

episodes are often associated with warm temperatures and stagnant meteorological conditions (NRC, 1991). O_3 concentrations in polluted atmosphere are affected by various meteorological parameters, such as temperature (Sillman and Samson, 1995), wind speed and vertical mixing (Sillman et al., 1995), and recirculation (Lu and Turco, 1996). Recent examples of the important links between meteorological parameters and the O_3 concentration are demonstrated in modeling studies by Baertsch-Ritter et al. (2004) and Dawson et al. (2007). The O_3 response to emissions of nitrogen oxides (NO_x) and volatile organic compounds (VOCs), critical information in formulating air quality management policy, is also affected by meteorology. For example, more stagnant meteorology (lighter wind speed and lower mixing height) tends to shift O_3 formation to the VOC-sensitive chemistry (Millford et al., 1994; Sillman et al., 1995; Sistla et al., 1996); temperature may have no direct effect on the O_3 sensitivity (Sillman, 1999), but higher temperature can change the size of the VOC-or NO_x -limited areas (Baertsch-Ritter et al., 2004). The effects of various meteorological parameters on the O_3 sensitivity chemistry in Europe have been substantiated recently by Baertsch-Ritter et al. (2004).

1 Introduction

The photochemistry in polluted troposphere, particularly the formation of ozone (O_3), is affected not only by the pollutant precursors and their chemical aging (Liu et al., 1987; Millford et al., 1994; Sillman, 1999), but also by meteorology (Solomon et al., 2000; Seaman 2000; Sillman, 1999 and references therein). It is well known that elevated O_3

As the most populous city in North America, the Mexico City Metropolitan Area (MCMA) exhibits degradation of air quality experienced by many megacities around the world (Molina and Molina, 2002, 2004). The intensive and extensive MCMA-2003 field campaign was conducted in April 2003 in the MCMA to improve our understanding of air pollution in Mexico City (Molina et al., 2007). The meteorological conditions during the campaign were identified and classified into three categories: “ O_3 -South”, “ O_3 -North” and “Cold Surge” by de Foy et al. (2005). These circulation patterns, shown in Fig. 1, were the results of the interaction of synoptic and local circulations. The “ O_3 -South” episode fea-



Correspondence to: L. T. Molina
(ltmolina@mit.edu)

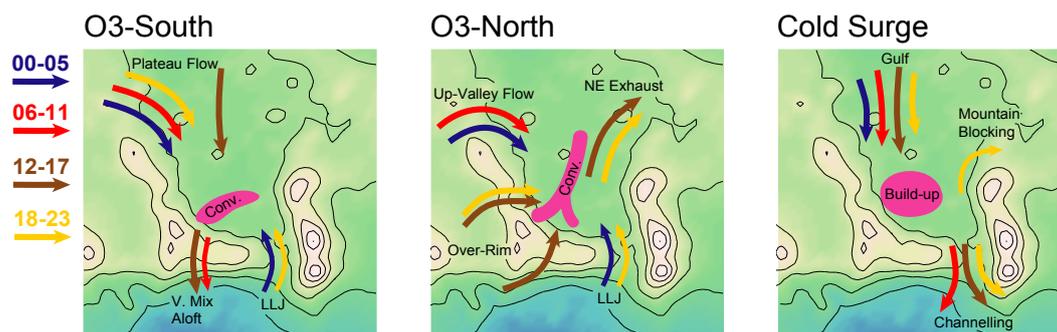


Fig. 1. Circulation model for O₃-South, Cold Surge and O₃-North episodes in the Mexico City basin. Arrows colored by time of day are dominant wind flows. Plateau winds on O₃-South days transport the urban plume southwards of the mountains. A low level jet forms in the Chalco passage causing a convergence line in the south of the city. Strong westerly flow on O₃-North days first enters the basin from the northwest but then comes over the south and west rims of the basin during the day. The combination with the low level jet forms a north-south convergence zone and exhaust of the basin to the northeast. On Cold Surge days strong northerly winds blow into the basin. Low mixing heights lead to pollutant accumulation in the city and channel flow through the Chalco passage.

tures a weak synoptic forcing associated with an anticyclone, and thermally driven circulations leads to high O₃ in the south of the city; the “O₃-North” episode features westerlies associated with weak anticyclone conditions to the south and a strong sub-tropical jet to the north of the MCMA, resulting in high O₃ in the north of the city; the “Cold Surge” episode characterizes cold northerlies from the Gulf that brings cold humid air to the Mexico plateau, leading to cloudy sky and afternoon convection and rainfalls in the MCMA and high O₃ in the city center due to stable conditions prior to the convection. These meteorological conditions also occur frequently in the MCMA in other years (de Foy et al., 2007).

Using a 3-D chemical transport model, CAMx 4.03 (Environ, 2003), we have previously examined the characteristics of photochemical O₃ formation and its response to precursors in the MCMA under the “O₃-South” episode (Lei et al., 2007). In this study, we extend the examinations to the other two meteorological events, “O₃-North” and “Cold Surge”, and compare with those obtained under the “O₃-South” episode. In addition, we have investigated the evolution of O₃ formation and response to the emissions reduction in the urban plume. One of our goals is to investigate whether and how these characteristics (such as the O₃ formation-precursor-radical relationship and the response of O₃ formation to changes in emissions) change under different meteorological conditions. If the NO_x/VOC chemistry of O₃ formation behaves consistently under different meteorological conditions, the formulation of emissions control strategies would be greatly facilitated. Another goal is to explore how the characteristics of O₃ formation and its sensitivity to NO_x and VOC levels evolve as the urban plume is transported downwind of the city, which has important implications for the regional air quality impacts of the urban plumes. Section 2 gives a brief description of the method. Section 3 presents the evaluation of the emission inventory, model performance, characteristics of O₃ formation and its

response to precursors, the chemical evolution of O₃ formation and its sensitivity to emissions reduction. A summary of conclusions is provided in Sect. 4.

2 Methodology

The 3-D chemical transport model CAMx 4.03 (Environ, 2003) with the SAPRC-99 gas phase chemical mechanism (Carter, 2000) was employed in this study. The model domain and configuration were the same as described in Lei et al. (2007). Briefly, the model domain covers 52×52 grids with a grid resolution of 3 km centered in Mexico City; there are 15 vertical layers extending from the surface to about 5 km a.g.l. with the bottom model layer interface of about 64 m a.g.l. The chemical initial and boundary conditions were the same as those used in Lei et al. (2007). Photolysis rate frequencies were pre-computed with the TUV model (Madronich and Flocke, 1998) for clear sky, and were corrected for the cloud effects using the approach of Chang et al. (1987). The meteorological inputs for the model were generated by MM5 (de Foy et al., 2006). The vertical diffusivity fields (k_v) were reconstructed from the state variables of the MM5 output using the CMAQ algorithm (Byun, 1999). de Foy et al. (2007) found that the CMAQ scheme overestimates k_v values. Accordingly we downscaled k_v to 20–30% based on comparison of modeled CO, NO_y concentrations and their ratios with observations.

The emission input fields were constructed by interpolating the official emissions inventories (EI) for the years 2002 and 2004 for the MCMA (CAM, 2004, 2006), and were adjusted based on measurements of CO, NO_y and speciated VOCs during the MCMA-2003 campaign (see Sect. 3.1).

In addition to the continuous measurements of a few critical pollutants by the monitoring stations of the Ambient Air Monitoring Network of Mexico City (RAMA) (SIMAT,

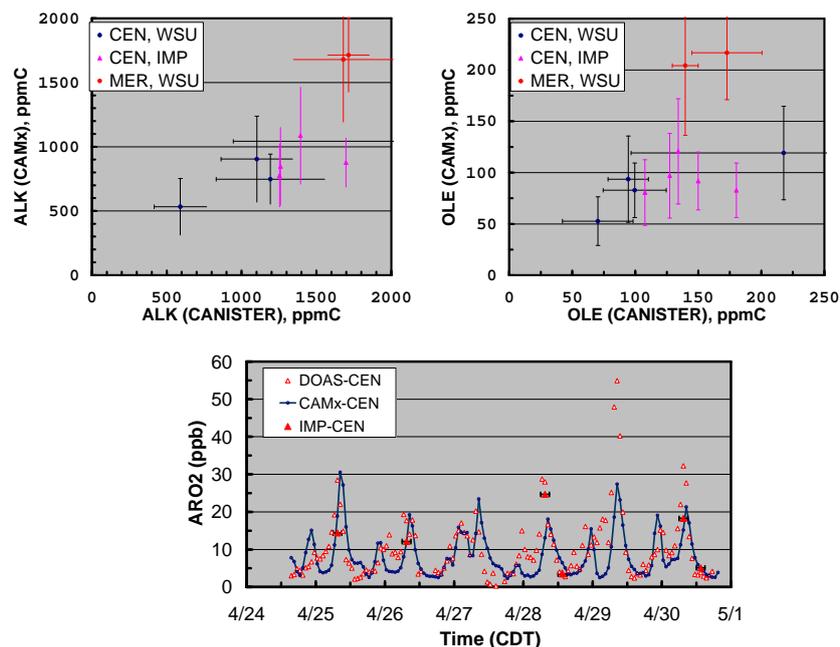


Fig. 2. Comparisons of modeled VOC concentrations with Canister (upper panel, 06:00–09:00 CDT average) and DOAS (lower panel) measurements after EI VOCs were adjusted. The error bar represents 1- σ standard deviation.

2008), an extensive array, in terms of time, species, and measurement technique, of near ground-level measurements for speciated VOCs were made during the MCMA-2003 campaign, which included speciated VOC measurements from canister sampling analyzed by Gas Chromatography/Flame Ionization Detection (GC/FID) at various sites (Velasco et al., 2007) and long-path Differential Optical Absorption Spectroscopy (DOAS) measurements of aromatics and formaldehyde at the campaign supersite CENICA (Volkamer et al., 2005). Evaluations of the model performance and the emission inventory were carried out against these ground-level measurements, i.e., CO, NO_y and O₃ were compared with the RAMA data, and speciated VOCs were compared with the GC-FID and DOAS datasets. The model performance and emission evaluation will be presented in the next section.

The simulation periods selected were 8–11 April 2003, which was characterized as a “Cold Surge” episode, and 25–30 April 2003, which was an “O₃-North” episode (de Foy et al., 2005). Simulations started one day earlier for each period allowing 24 h for spin-up.

3 Results and discussions

3.1 Evaluation of emission inventory

Lei et al. (2007) evaluated the official EI for the year of 2002 using the RAMA data, VOC dataset from the canister sampling coupled with the GC/FID analysis, and the DOAS measurements at CENICA during 13–15 April 2003 (the

“O₃-South” episode). In this work, we extended the evaluation of EIs for the years 2002 and 2004 using a VOC dataset with a broader spatial (three sites) and temporal coverage (8–11 and 25–30 April). The GC/FID measurements used in this study included data obtained at CENICA, Pedregal and La Merced during the MCMA-2003 by research groups from Washington State University (WSU) and the Mexican Petroleum Institute (IMP) (Velasco et al., 2007); the aromatics and HCHO concentrations measured with DOAS covered the selected simulation periods at CENICA. The evaluation procedure is described in detail in Lei et al. (2007). In brief, first, a spatially and temporally resolved and chemically speciated initial emission estimate was constructed from the annual emissions in the official emission inventory; second, the initial emission estimate was then adjusted based on the RAMA observations of CO and NO_y and the MCMA-2003 field measurements of speciated VOCs. Model runs with varying emission scaling factors for each gaseous primary model species were repeated and simulated, and morning rush hour (06:00–09:00 a.m.) concentrations of CO, NO_y and speciated VOCs were compared with measurements until a good model-measurement agreement for these species was reached. The comparisons were carried out for all primary model VOC species except higher aldehydes (CCHO and RCHO) whose measurements were unavailable or incomplete.

Figure 2 gives an example of the VOC comparisons after the emissions were adjusted, which shows a good agreement between simulated and measured VOCs. Table 1 summarizes the emission adjustment factors, which shows

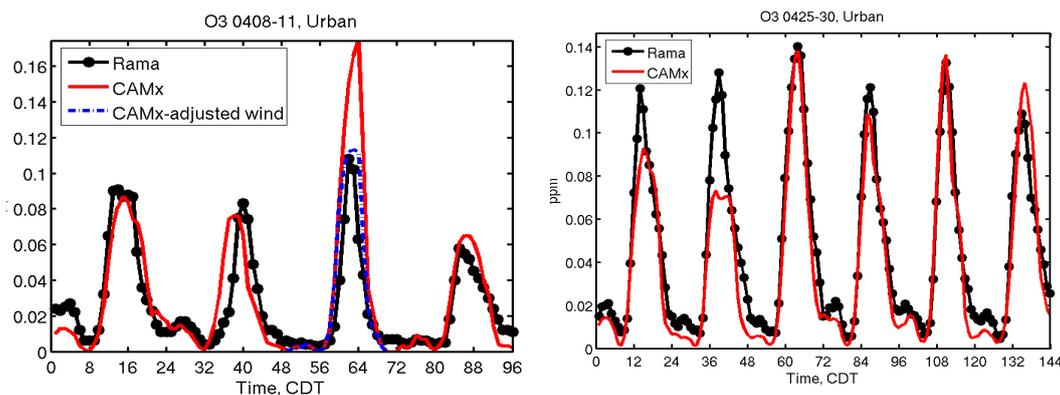


Fig. 3. Comparison of measured (black) and simulated (red) diurnal variation of near surface O_3 concentrations averaged over the urban area (21 monitoring stations) during the “Cold Surge” (8–11 April, left) and the “ O_3 -North” (25–30 April, right) episodes. Also shown in the left panel are the simulated O_3 concentrations (blue dashed) on 10 April when the gap flow was enhanced by 0.5 m/s.

Table 1. Adjustment factors for correction of the MCMA emissions inventory used in this study.

CAMx Model species	Adjustment factor
CO	1.0
NO_x	1.0
ALK1	2.5
ALK2	3.0
ALK3	4.0
ALK4	1.2
ALK5	0.6
ETHE	1.2
OLE1	0.9
OLE2	0.8
ARO1	1.2
ARO2	1.2
HCHO	7.0
CCHO and RCHO	4.0

that the emission estimates for CO and NO_x in the official EIs are generally accurate; the emissions of most alkanes are underestimated by about a factor of 2–4 while the most active alkanes (C7 or above, denoted as ALK5 in Table 1) are overestimated by about a factor of 2; emissions of alkenes are generally accurate, and the emissions of aromatics are probably slightly underestimated by $\sim 15\%$. The overestimate of alkanes of C7 or above is probably an artifact due to their low concentrations, which may impose a larger uncertainty in the measurement, and also because some significant species may not be detected in the measurements. The serious underestimate of HCHO emissions in the MCMA are consistent with the findings of Kolb et al. (2004) and Zavala et al. (2006). An adjustment factor of 4 was assumed for higher aldehydes. The adjustment factors obtained in this study are consistent with those reported in Lei et al. (2007),

given the difference in different EI years, locations where measurements were made, and the size of dataset used for comparison. The adjusted emissions were then used for the base case run.

It should be noted that even though we have broadened the dataset used for the EI evaluations, the VOC comparisons were still made over limited number of sites (CENICA, Pedregal and La Merced) and over a relatively short time period (a few weeks). A more complete EI evaluation using this approach requires more measurements with broader spatial and temporal coverage. In addition, direct evaluations of emissions using other measurement techniques (e.g., Kolb et al., 2004; Zavala et al., 2006; Velasco et al., 2007) are needed.

3.2 Performance of O_3 simulation

Figure 3 shows the comparison of measured and simulated near-surface O_3 concentrations in the MCMA urban region during 8–11 (“Cold Surge”) and 25–30 (“ O_3 -North”) April 2003. Overall the model reproduced the observations quite well except on some days, such as April 10 and 26. Figure 4 shows that on 10 April the simulated O_3 was too high, perhaps because the simulated northerly component (gap flow) was too weak, leading to the over accumulation of air pollutant plume inside the basin. On 26 April, the simulated photochemical plume in the basin was shifted to the north due to overestimated SW flow (channel flow) in the morning hours, deviating far from the plume’s average location and transport pathway during this episode.

A sensitivity test with an uniform increase of the gap flow by 0.5 m/s (both the northerly and westerly components) between 10:00–17:00 CDT on 10 April dramatically improved the O_3 agreement (Fig. 3); a test with a weakening of the SW channel flow on 26 April (not shown) also significantly improved the O_3 agreement. de Foy et al. (2006) find that the MM5 model bias for April 2003 in the MCMA is between -0.5 – 0.0 m/s for surface winds and the variance of

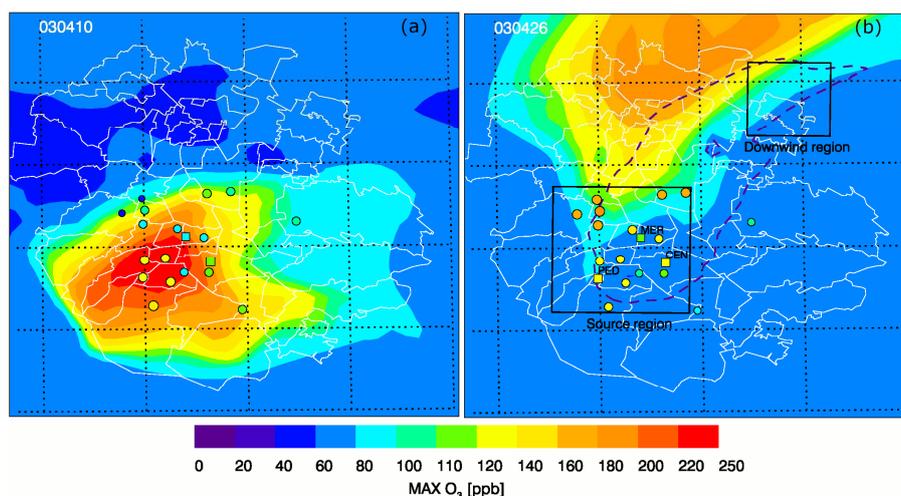


Fig. 4. Comparison of simulated (colored contour) versus observed (colored dots and squares) geographical distribution of near-surface peak O_3 concentration on (a) 10 April and (b) 26 April 2003 in which the model did not perform well. The area encompassed by the dashed line in (b) is the area where peak ozone averaged over 25–30 April (except 26) 2003 exceeded 120 ppb, reflecting the plume’s average location and pathway during the “ O_3 -North” episode. The near source area and the downwind area where the data are sampled are also delimited in (b). Also shown are the three VOC canister sampling sites.

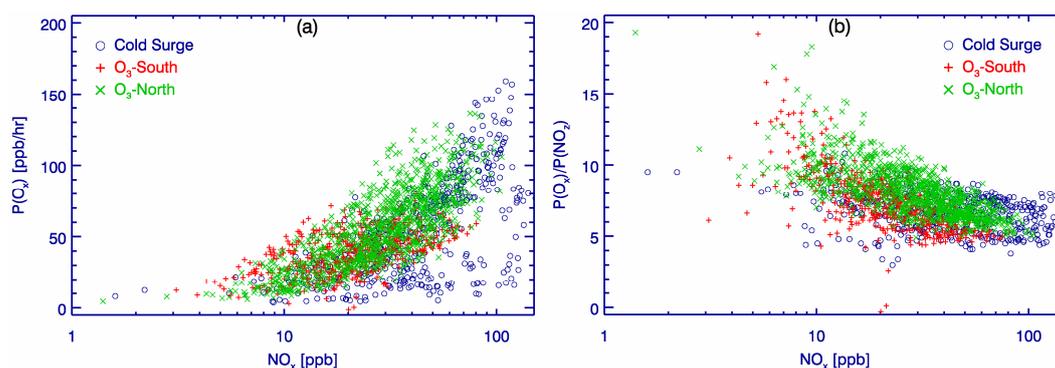


Fig. 5. Simulated midday (12:00–17:00 CDT) net O_x production rate (left) and O_3 production efficiency (right) as a function of NO_x concentration in the urban area of MCMA under different meteorological conditions. Data points are reduced by using every 3rd data point to avoid the high density of points.

errors is 1.0–1.5 m/s (c.f. Fig. 5 in de Foy et al.); the model bias for the upper level winds is similar (c.f. Fig. 18 in de Foy et al.). This suggests that the wind speed modifications used in the above tests are within the model uncertainty, and it also suggests that there is room for improving the meteorological model. These results indicate that O_3 concentration on the urban scale is very sensitive to the wind field, and it is essential that the chemical transport model be driven by an accurate meteorological input in the urban scale photochemical modeling. Due to the meteorological and chemical discrepancies, simulated data for 10 and 26 April will not be included for further discussion.

Table 2 summarizes the statistical performance for major gaseous pollutants in the MCMA (over 21 RAMA monitoring stations in the urban setting) during each episode (with

the exclusion of 10 and 26 April). It should be noted that during the “Cold Surge” episode the surface O_3 concentration was lower compared to other meteorological conditions. This is because there was a stronger daytime horizontal advection on the “Cold Surge” days. In contrast, CO and NO_y concentrations during the Cold Surge episode were similar to those during the “ O_3 -North” episode because the mean values of CO and NO_y were weighted more by the concentrations at morning hours and night-time when the surface wind was generally calm for all episodes. During the “ O_3 -South” episode (13–15 April 2003), the O_3 concentration was similar to that of the “ O_3 -North” episode, but the primary pollutant (CO and NO_y) concentrations were lower than those during other meteorological conditions, which is probably due to reduced emissions during this episode. The

Table 2. Model statistical performance measures in the MCMA.

			O ₃ *	CO	NO _y
Cold Surge	Mean obs	(ppb)	27.1	1887	80.9
	Mean mdl	(ppb)	26.4	1642	75.5
	RMSE	(ppb)	18.9	1313	55.3
	IOA		0.88	0.78	0.80
	NB	(%)	-5.6	49.4	33.4
	NGE	(%)	30.6	86.5	66.6
	NMB	(%)	-9.5	-13.1	-6.3
O ₃ South	AAP	(%)	7.5	-11.6	-4.3
	Mean obs	(ppb)	47.5	1273	61.0
	Mean mdl	(ppb)	42.7	1379	60.6
	RMSE	(ppb)	21.6	925	43.0
	IOA		0.93	0.76	0.79
	NB	(%)	-1.7	71.5	30.3
	NGE	(%)	24.6	97.2	63.8
O ₃ North	NMB	(%)	-1.3	8.2	-1.1
	AAP	(%)	8.9	28.6	10.1
	Mean obs	(ppb)	45.8	1884	84.2
	Mean mdl	(ppb)	41.0	1894	83.8
	RMSE	(ppb)	24.9	1277	58.3
	IOA		0.92	0.77	0.81
	NB	(%)	-5.6	47.9	36.4
	NGE	(%)	28.3	79.3	68.4
	NMB	(%)	-5.9	0.6	-0.4
	AAP	(%)	3.2	15.0	2.0

Obs = observation, Mdl = model, RMSE = root mean square error, IOA = index of agreement (Willmott, 1981), NB = normalized bias, NGE = normalized gross error, NMB = normalized mean bias, AAP = average accuracy of peak paired in site.

*: in the calculation of NB, NGE and NMB for O₃, the threshold value for observed O₃ was set to 40 ppb.

episode of 13–15 April 2003 started on Sunday followed by the Easter Week; anthropogenic emissions were reduced by 10–20% during this episode (Lei et al., 2007) when many residents left the city for holidays. The result of similar O₃ but less CO and NO_y on weekends compared to weekdays is in agreement with the finding of a recent study of the weekend effect in the MCMA based on long-term RAMA measurements (Stephens et al., 2008).

3.3 Characteristics of O₃ production rate

We examined the simulated net photochemical formation rates of O_x (O_x=O₃+NO₂), P(O_x), and ozone production efficiency (OPE), as a function of NO_x during 8–11, 13–15, and 25–30 April; the results are presented in Fig. 5. For all three meteorological categories, generally P(O_x) increases with increasing NO_x despite various degree of scatter. P(O_x) values are higher during the “O₃-North” and “Cold Surge”

episodes than those during the “O₃-South” episode. The lower P(O_x) values in the “O₃-South” episode are due to the reduced emissions as mentioned above, which leads to decreased VOC reactivity and radical production that will be illustrated later. The relatively high values of P(O_x) during the “Cold Surge” episode are probably in part due to the fact that MM5 may not completely capture the cloud effects that occurred frequently under this type of meteorological condition, which affect the air mixing (de Foy et al., 2007) and photochemical activities. Despite the high P(O_x) values, the O₃ concentrations during the “Cold Surge” episode are lower than other meteorological conditions (see Table 2 and Fig. 3) due to the stronger advection during this event. During the “Cold Surge” event, there are many data points with high NO_x but low P(O_x), which may be due to the imminent impacts of emissions at some locations (see Fig. 6b below). The values of OPE (concentrating at 5–10 molecules of O₃ per NO_x oxidized) and its dependence on NO_x are similar and consistent under “O₃-South”, “O₃-North” and “Cold Surge” conditions, although under the “Cold Surge” condition OPE values appear to be less NO_x-dependent. In summary, the characteristics of P(O_x) and OPE are consistent under all meteorological conditions.

We have analyzed the relationship among P(O_x), VOC reactivity (K_{VOC}), chemical aging (NO_z/NO_y), and primary radical production rate (Q) in the urban area under different meteorological conditions; the results are shown in Fig. 6. The relationship between P(O_x) and K_{VOC} (Fig. 6a) is similar to that of P(O_x)-NO_x, indicating the co-emissions of NO_x and VOCs. Figure 6b shows that in the “Cold Surge” event, the urban plume is fresher, with less chemical processing, which results in high NO_x and VOCs, but low P(O_x). In the less aged air with high NO_x and VOCs, the O₃ concentration is lower because of titration by NO and radical scavenging by NO_x, leading to lower radical sources from O₃ photolysis and other radical precursors (such as aldehyde photolysis and alkene ozonolysis) due to slower photochemical processes. Since P(O_x) is strongly dependent on the radical sources (Fig. 6c), lower Q values lead to lower P(O_x). The urban air mass is most chemically aged during the “O₃-South” episode, which is understandable considering the dominant flow patterns and terrain setting. As will be discussed later, chemical aging has significant influences on the NO_x-VOC sensitivity of O₃ formation. Over 90% of the radicals are removed through the radical-NO_x reactions (Fig. 6d) compared to less than 10% via the radical-radical reactions, implying that the midday O₃ formation is VOC-limited (Sillman, 1995; Kleinman et al., 1997; Daum et al., 2000). Q and K_{VOC} values (which are closely related to each other) are higher during the “Cold Surge” and “O₃-North” episodes than during the “O₃-South” episode, leading to higher P(O_x) values in the former. The lower Q and K_{VOC} values during the “O₃-South” episode are probably due to reduced emissions.

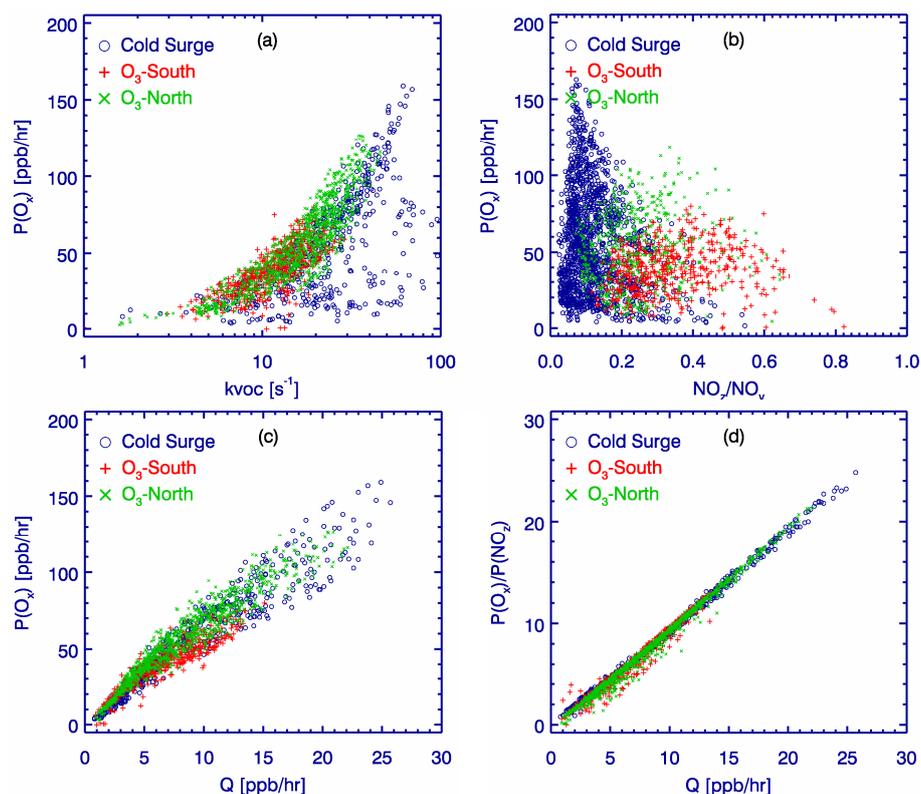


Fig. 6. Simulated relationship among $P(\text{O}_3)$, VOC reactivity (k_{VOC}), chemical aging (NO_z/NO_y) in the urban area between 12:00–17:00 CDT during different meteorological conditions. Data points are reduced by using every 3rd data point to avoid the high density of points.

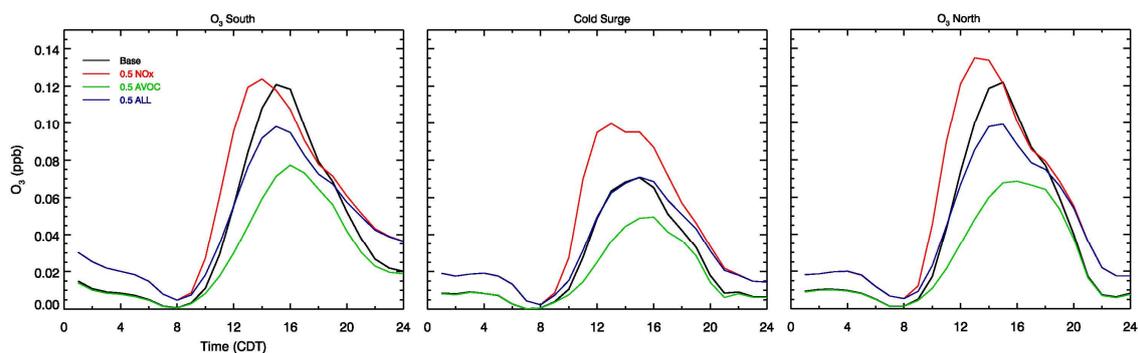


Fig. 7. Time series of surface O_3 concentrations for different emission control scenarios under different meteorological conditions. Data are averaged over 8 urban representative monitoring stations and over each episode. In the legend Base denotes the base case, 0.5NO denotes a 50% reduction in NO_x emissions, 0.5VOC a 50% reduction in VOC emissions, and 0.5ALL a 50% reduction in both NO_x and VOC emissions.

3.4 Response of O_3 production to precursors

We have evaluated the response of O_3 formation to the O_3 precursors by perturbing the emissions and examining the resulting changes in O_3 concentrations and O_3 production rates. The perturbations of emissions include a 50% reduction in NO_x emissions only, a 50% reduction in VOC emissions only, and a 50% reduction in both NO_x and VOC emissions.

Figure 7 illustrates the effects of emission reductions on surface O_3 in the urban region. It shows that a 50% reduction in NO_x emissions leads to an increase in O_3 , while a 50% reduction in VOC emissions leads to a decrease in O_3 by about a factor of 2. In the NO_x 50% reduction scenario, O_3 increases marginally during the “ O_3 -South” episode, but clearly increases during the other two episodes, most notably

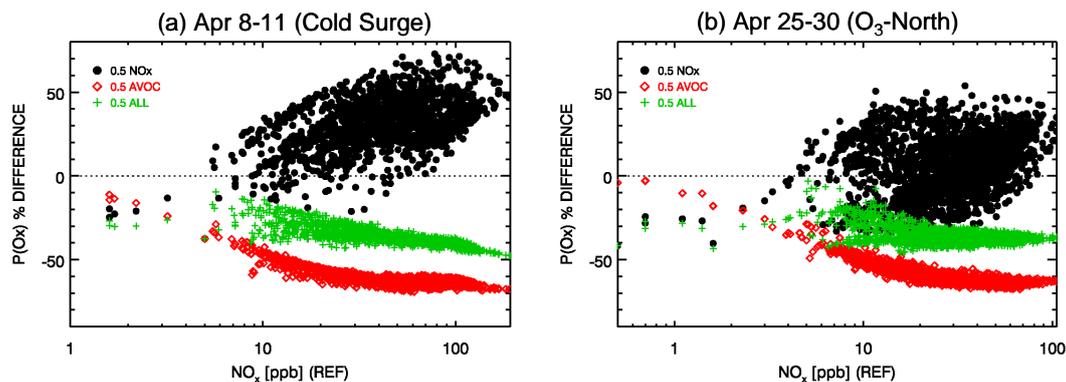


Fig. 8. The percentage change of $P(O_x)$ as a function of NO_x from 12:00–17:00 CDT during (a) the “Cold Surge” and (b) the “ O_3 -North” episodes in the MCMA urban region.

during the “Cold Surge” episode. This response behavior is consistent with the chemical aging of the air mass: least aged during the “Cold Surge” days, most aged during the “ O_3 -South” days, and the “ O_3 -North” days in between. Another interesting phenomenon occurs in the 50% reduction in both NO_x and VOC emissions scenario. During the “ O_3 -North” episode, the surface O_3 concentrations decrease as in the “ O_3 -South” case; however, during the “Cold Surge” episode, the midday O_3 does not change noticeably. This does not mean that O_3 sensitivity chemistry changes during the “Cold Surge” condition compared to other meteorological conditions, as will be discussed below.

Figure 8 shows the percentage change of $P(O_x)$ as a function of NO_x from 12:00–17:00 CDT during the “Cold Surge” and “North-north” episodes. As in the “ O_3 -South” case, a 50% reduction in VOCs leads to a significant decrease in $P(O_x)$; a 50% reduction in NO_x leads to a decrease in $P(O_x)$ at low NO_x conditions but an increase in $P(O_x)$ at high NO_x ; the 50% reduction in both NO_x and VOC also leads to a decrease in $P(O_x)$ but less than that of the VOC-only reduction case. This is different from the change in O_3 concentration showed in Fig. 7, because $P(O_x)$ is more dependent on chemical cycles while O_3 concentration is sensitive to both chemistry and meteorology. The inconsistency in the behaviors of O_3 concentration and $P(O_x)$ suggests that although meteorology may affect O_3 concentrations significantly, it may not affect the rate of O_3 photochemical production as much. The response of $P(O_x)$ to emission reductions indicates that midday O_3 formation is VOC limited in the urban region, independent of meteorological conditions.

3.5 Evolution of photochemical plume

In order to explore the evolution of a pollution plume as it transports pollutants downwind of the city and its impacts on the regional pollutant levels, we examined the evolution of $P(O_x)$ in a pollutant plume following its pathway during the “ O_3 -North” event. The “ O_3 -North” event was chosen

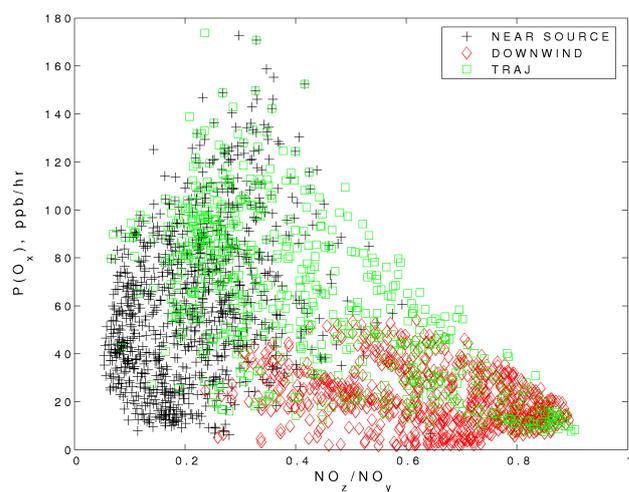


Fig. 9. Relationship between $P(O_x)$ and chemical aging (NO_z/NO_y) at 12:00–17:00 CDT during 25–30 April 2003. Black dots are data sampled in the downtown near the source area, red diamonds are data sampled about 60 km downwind of the urban core, and green squares are data sampled along the surface plume trajectories. Data points in the near source and the downwind areas are reduced by using every 2nd data point to avoid the high density of points.

because during this event the MCMA urban plume travels over a larger space, exerting most environmental impacts at regional scales. Figure 9 shows the evolution of $P(O_x)$ in a plume with three different origins: source area, downwind area, and along the transport pathway. The source and downwind area where the data were sampled are indicated in Fig. 4b. The transport path was identified by visually tracking the location of relatively high surface CO concentrations coincident with high O_3 concentrations each hour during 12:00–17:00 CDT every day. We observe that there are distinct relationships between $P(O_x)$ and chemical aging (using NO_z/NO_y as the indicator for the photochemical

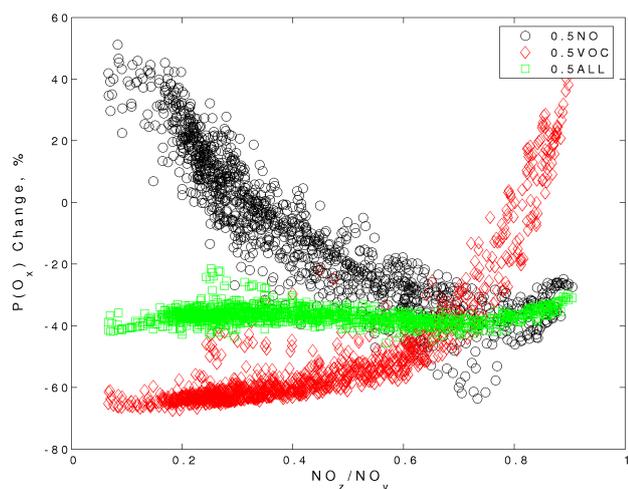


Fig. 10. Percentage change of $P(O_x)$ as a function of chemical aging along the plume transport pathway at 12:00–17:00 CDT during 25–30 April 2003 when emissions of NO_x and VOCs are reduced by 50% sequentially and simultaneously.

aging) for plumes in the urban center and in the downwind areas. Plumes in the source area are much fresher (indicated by low NO_z/NO_y values); as the plumes undergo chemical processing, $P(O_x)$ increases rapidly and reaches very high values within a narrow range of the chemical age parameter. On the other hand, plumes in the downwind area, about 60 km from the source area, are mostly highly chemically aged, and $P(O_x)$ values are low and nearly constant. Plumes along the transport pathway are characterized by a combination of the signatures of fresh plumes (source area) and aged plumes (downwind), this is because as the plume is transported downwind, it becomes progressively chemically aged, even as continuous emissions along the pathway are mixed into the plume.

We also investigated the evolution of O_3 formation sensitivity along the plume's transport pathway. Figure 10 shows the percentage change in $P(O_x)$ as a function of chemical aging when emissions are reduced (50% in NO_x only, 50% in VOCs only, and 50% both) during 12:00–17:00 CDT, 25–30 April 2003. It can be seen that the O_3 sensitivity changes as the plume evolves. When the plume is fresh, O_3 formation is more VOC sensitive (NO_x depressed); as the plume becomes chemically aged, O_3 formation becomes progressively VOC insensitive and more NO_x sensitive. When the ratio of NO_z/NO_y reaches 0.8 or higher, O_3 formation becomes NO_x sensitive. When the plume is highly aged, the VOC emission reduction tends to increase $P(O_x)$. This is probably because reducing VOCs would lead to lower radical concentrations and thus lower NO_z formation (through the radical- NO_x reactions), which would result in higher production of NO_2 , even though $P(O_3)$ changes little ($O_x=O_3+NO_2$).

4 Conclusions

We have extended the characterization of O_3 formation and its response to emission reductions under “ O_3 -South” (13–15 April 2003) meteorological condition to two more distinct episodes (8–11 April “Cold Surge” and 25–30 April “ O_3 -North”) using the CAMx model with an improved emission estimate evaluated with greater spatial and temporal coverage. Further, we have examined the evolution of the O_3 formation and its response to emissions reduction using the “ O_3 -North” episode as a case scenario.

The accuracy of the MCMA EI was re-evaluated using observed dataset (particularly speciated VOCs) with greater spatial and temporal coverage. The results show that the estimates of CO and NO_x emissions in the MCMA EI are accurate, while VOCs are in general underestimated by a factor of 2–3 (1.6 overall), consistent with the study of Lei et al. (2007). However, even though a larger dataset was used for the EI evaluations, the VOC comparisons were still made over limited locations (CENICA, Pedregal and La Merced) and over a relatively short time period (a few weeks). A thorough evaluation of the EI accuracy requires more measurements with a broader spatial and temporal coverage, and measurements from other techniques such as the fleet chasing technique.

By analyzing the relationship among O_x production rate $P(O_x)$, O_3 precursors and radical sources, we found that the characteristics of O_3 formation and its response to emission changes are weakly dependent on meteorological conditions. Under all three meteorological categories, in general midday $P(O_x)$ increases with NO_x with no obvious turnaround behavior, and both midday $P(O_x)$ and ozone production efficiency are high at high NO_x conditions compared to those in other urban cities; O_3 formation is VOC limited in the MCMA urban source region as revealed by both the dependence behavior of $P(O_x)$ on NO_x and VOCs levels and radical production rate and the response of $P(O_x)$ to precursor emission reductions.

We have also examined the evolution of O_3 production and its response to emission reductions during the “ O_3 -North” event. The plumes in the source region are found to be less chemically aged, and the O_3 production increases rapidly and reaches very high values within a short chemical age, while plumes in the downwind area are chemically aged and characterized by low O_3 production; plumes in between are featured by the combination of fresh and aged plumes. In the fresh plume O_3 formation is more VOC sensitive. As the plume becomes chemically aged, O_3 formation becomes progressively VOC insensitive and more NO_x sensitive.

It should be pointed out that in this study, we have covered less-than-a-week simulations for each of the three meteorological episodes. More studies are needed to comprehensively characterize the O_3 formation and its response to precursors under various meteorological conditions. We plan further modeling studies using the much larger and

comprehensive datasets from the 2006 MILAGRO (Megacity Initiative: Local and Global Research Observations) Campaign (Molina et al., 2008), in which the meteorology was characterized by 6 episode types representing different wind transport patterns (de Foy et al., 2008), in contrast to the three episode types for MCMA-2003, and where aircraft measurements provide unique opportunities for further model evaluations and interpretations and scientific explorations.

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